

Exploring the Mechanisms of Land-Atmosphere Interactions during HI-SCALE

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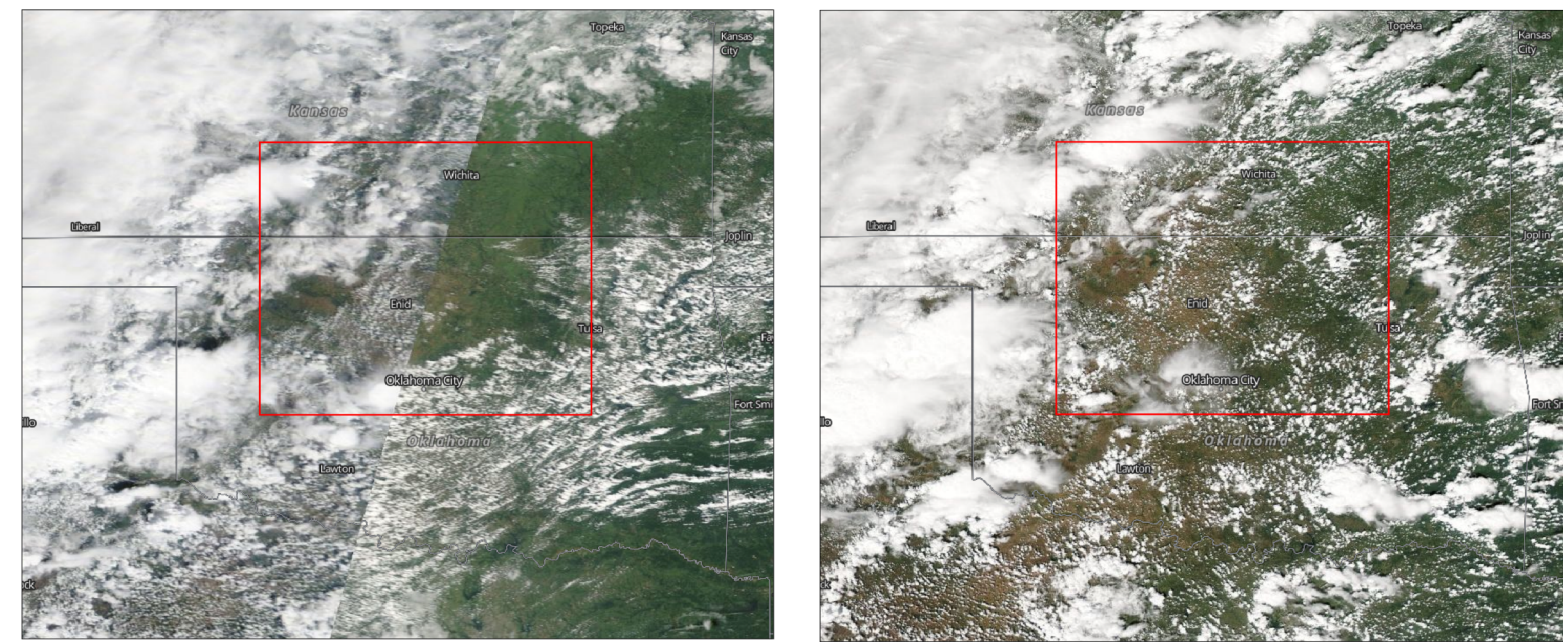
Introduction

Land-atmosphere interactions play important roles in the initiation of shallow convection and the subsequent potential transitions to deep convection. This study explores how the land properties are related to the convection and cloud populations during daytime over Southern Great Plain (SGP) based on an LES-version of WRF Simulations. Isentropic and cluster analysis of **Equivalent potential temperature (θ_e)** are used to assess how land surface conditions lead to the simulated cloud populations.

Case Overview

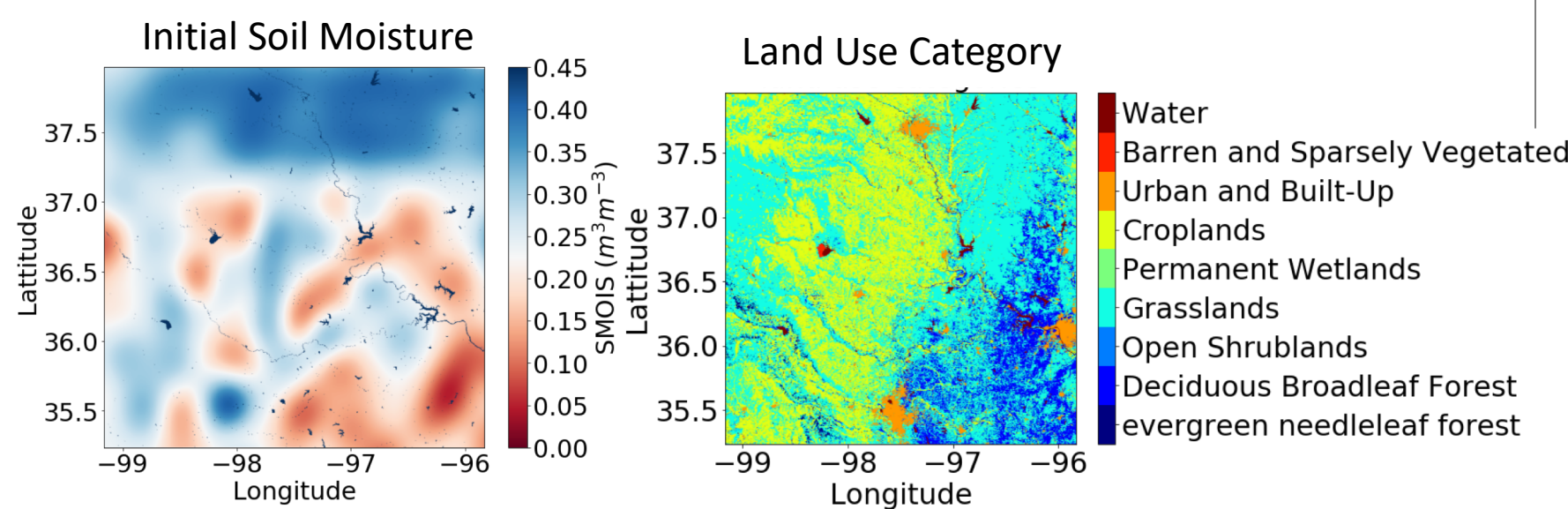
- Time and Location:** August 30th, 2016, Southern Great Plain
- A “Golden Day” with transitions from shallow to deep convection

MODIS Terra ~1630 UTC (~1030 CST) MODIS Aqua ~1950 UTC (~1350 CST)



WRF-LES Simulations

- From 08/30 12:00 UTC to 08/31 00:00 UTC
- Spatial Resolution: 300m
- Output time step: 15 min
- More details are in Jerome Fast’s talk (A41D03) on Thursday.



Isentropic Analysis

Isentropic analysis helps to

- reduce the 4D dataset into 3D (horizontal 2D $\rightarrow \theta_e$),
- separate the air between the ascent, warm, moist air and subsiding, cold, dry air, and
- isolate the irreversible convection overturning by filtering out reversible motions.

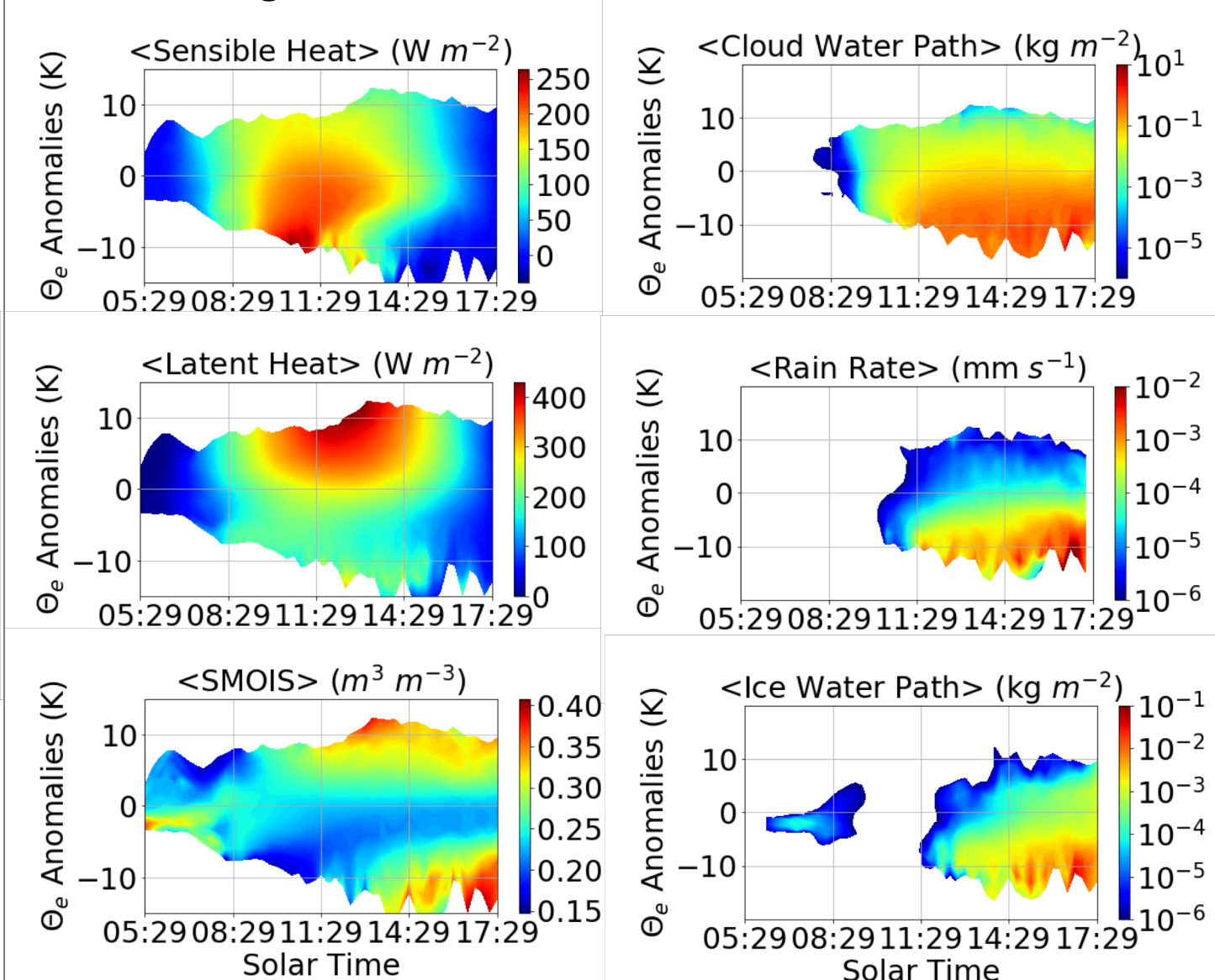


Figure 2. Isentropic distribution of surface properties.

Over surface (Figure 2): lowest level θ_e' (θ_{es}')

- High θ_{es}' are associated with high latent heat (LH) and low sensible heat (HFX).
- Soil moisture (SMOIS):
 - High initial SMOIS causes high LH, so the θ_{es}' of those region increase.
 - Raining events cause another high SMOIS region with low θ_{es}' .
- Regions with large cloud water path are over low θ_{es}' .

Land use type (Figure 3):

- Grassland and broadleaf forest are associated with high values of θ_{es}' , while cropland, urban and water are associated with low values of θ_{es}' .

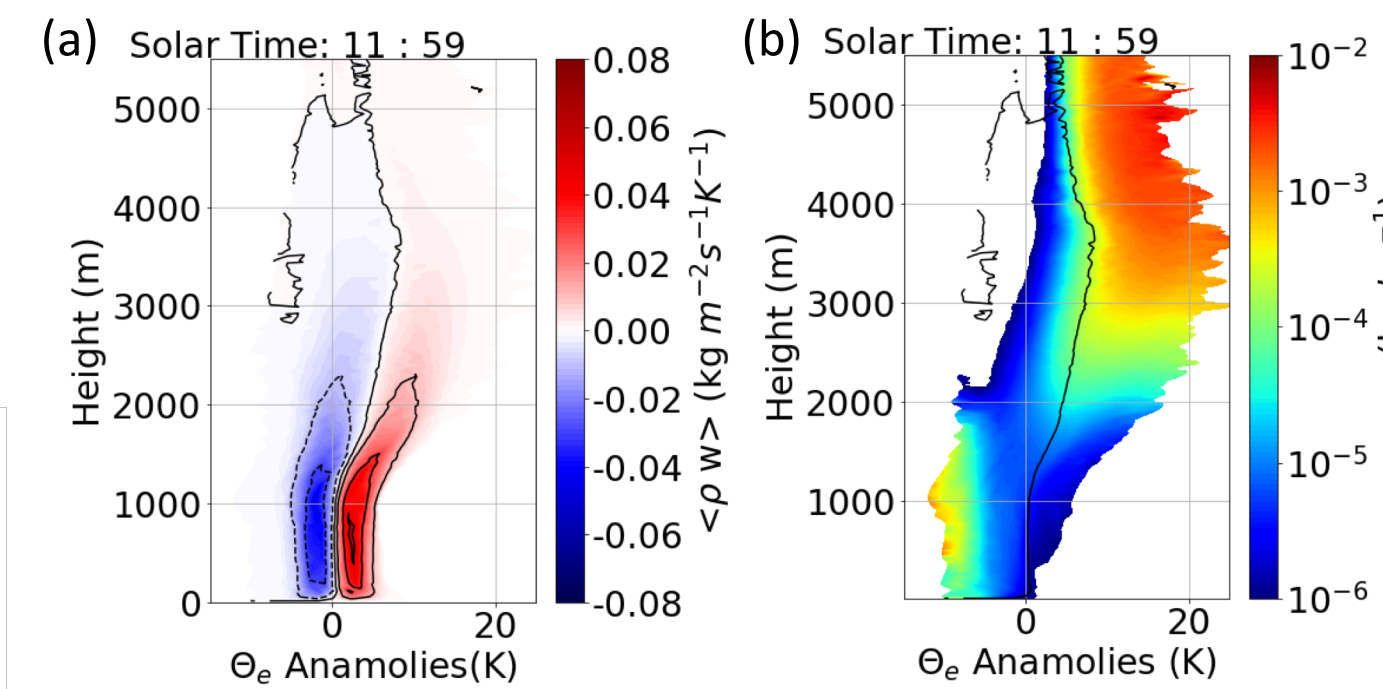


Figure 1. Isentropic distribution of (a) mass flux and (b) rain water mixing ratio.

$$\langle \rho x \rangle (z, \theta_{e0}) = \frac{1}{L_x L_y} \iint_{0.0}^{L_x L_y} [\rho x \delta(\theta_e - \theta_{e0})] dx dy \quad (\text{Eq.1})$$

$$\langle x \rangle = \langle \rho x \rangle / \langle \rho \rangle \quad (\text{Eq.2})$$

In lower troposphere (Figure 1):

- Upward mass flux is associated with positive θ_e anomalies (θ_e'), and vice versa.
- Rainwater starts to appear over the region with positive θ_e' , and falls on the region with negative θ_e' .

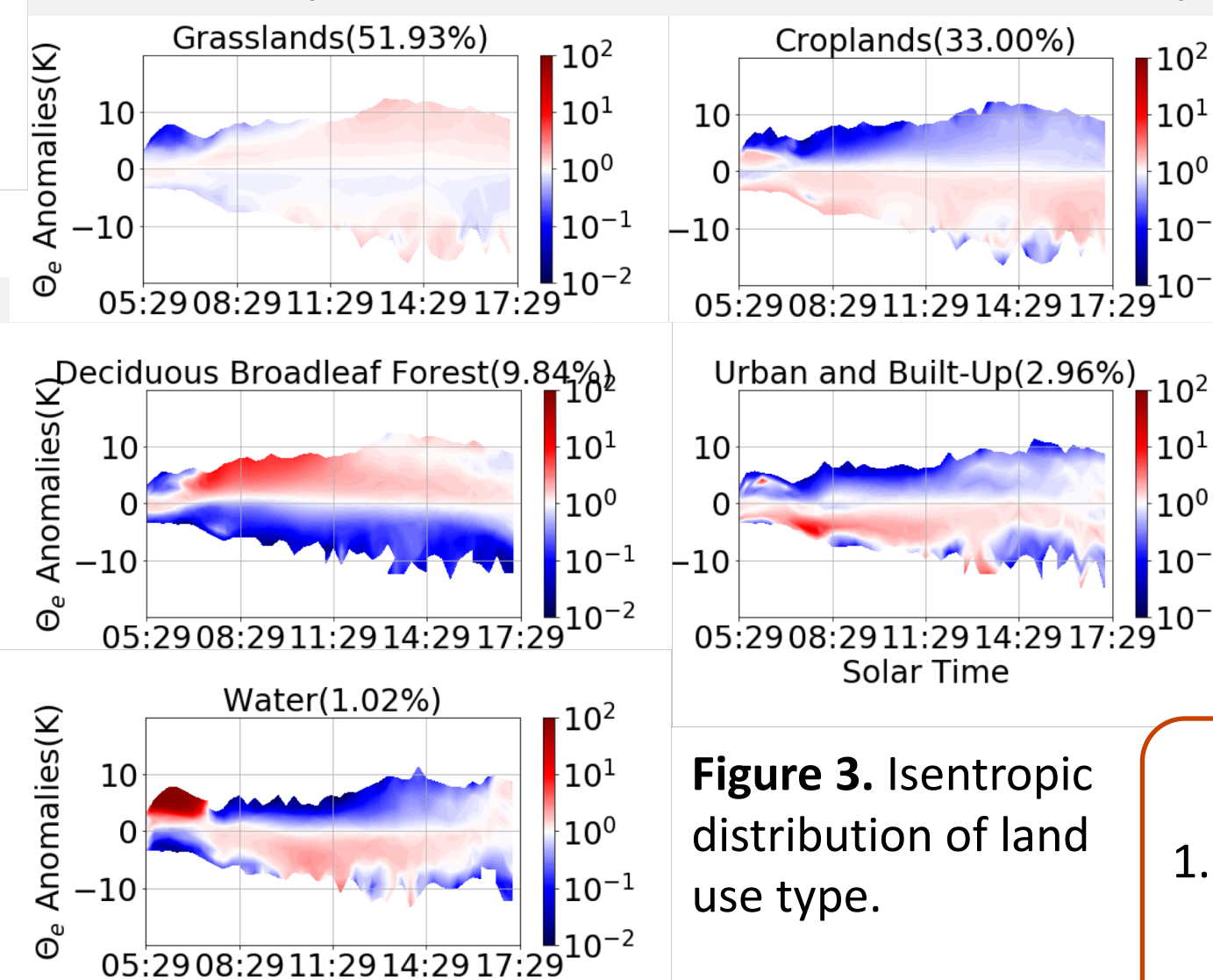


Figure 3. Isentropic distribution of land use type.

$$\frac{\langle LU_i \rangle (z, \theta_{e0})}{f_i} = \frac{\iint \delta(LU, LU_i) \delta(\theta_e, \theta_{e0}) dx dy}{\iint \delta(\theta_e, \theta_{e0}) dx dy} = \frac{1}{L_x L_y} \iint \delta(LU, LU_i) dx dy \quad (\text{Eq. 3})$$

Cluster Analysis

“K-means” unsupervised learning (Figure 4)

- Samples: θ_{es}' over all the grids (990×990)
- Features: 49 time steps
- Eight clusters
- (Figure 4b) The trends of θ_{es}' time series among clusters are dramatically different.

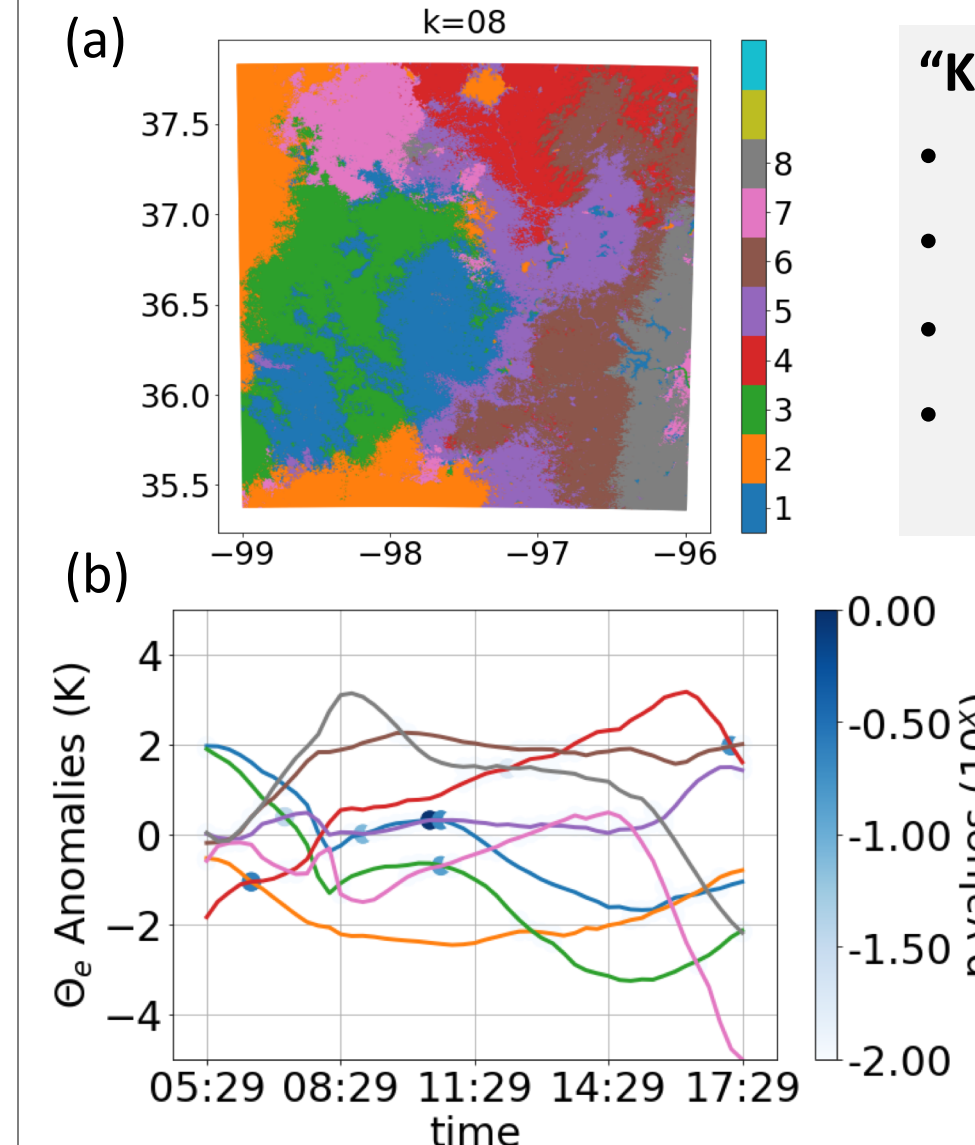


Figure 4. (a) Cluster analysis of θ_{es}' . (b) θ_{es}' time series of each cluster.

Budgets of θ_e (Figure 5):

$$\ln \theta_e = \ln(\theta) + c_1 q_v \quad (c_1 \text{ is constant})$$

$$\rightarrow d \ln \theta_e / dt = d \ln \theta / dt + c_1 dq_v / dt$$

- Contributions from $d \ln \theta / dt$ to $d \ln \theta_e / dt$ are more than those from $c_1 dq_v / dt$.
- Most of the variations among clusters are explained by $c_1 dq_v / dt$.

- Period 1 (5:29-8:29): High LH causes high $c_1 dq_v / dt$, hence high $d \ln \theta_e / dt$ and increasing θ_{es}' .
- Period 2 (8:29-11:29): Negative $c_1 dq_v / dt$ and positive $d \ln \theta / dt$ are balanced.
- Period 3 (11:29-14:29): Rain processes compensate the water vapor depletion.
- Period 4 (14:29-17:29): Similar with P3 except for more rain.

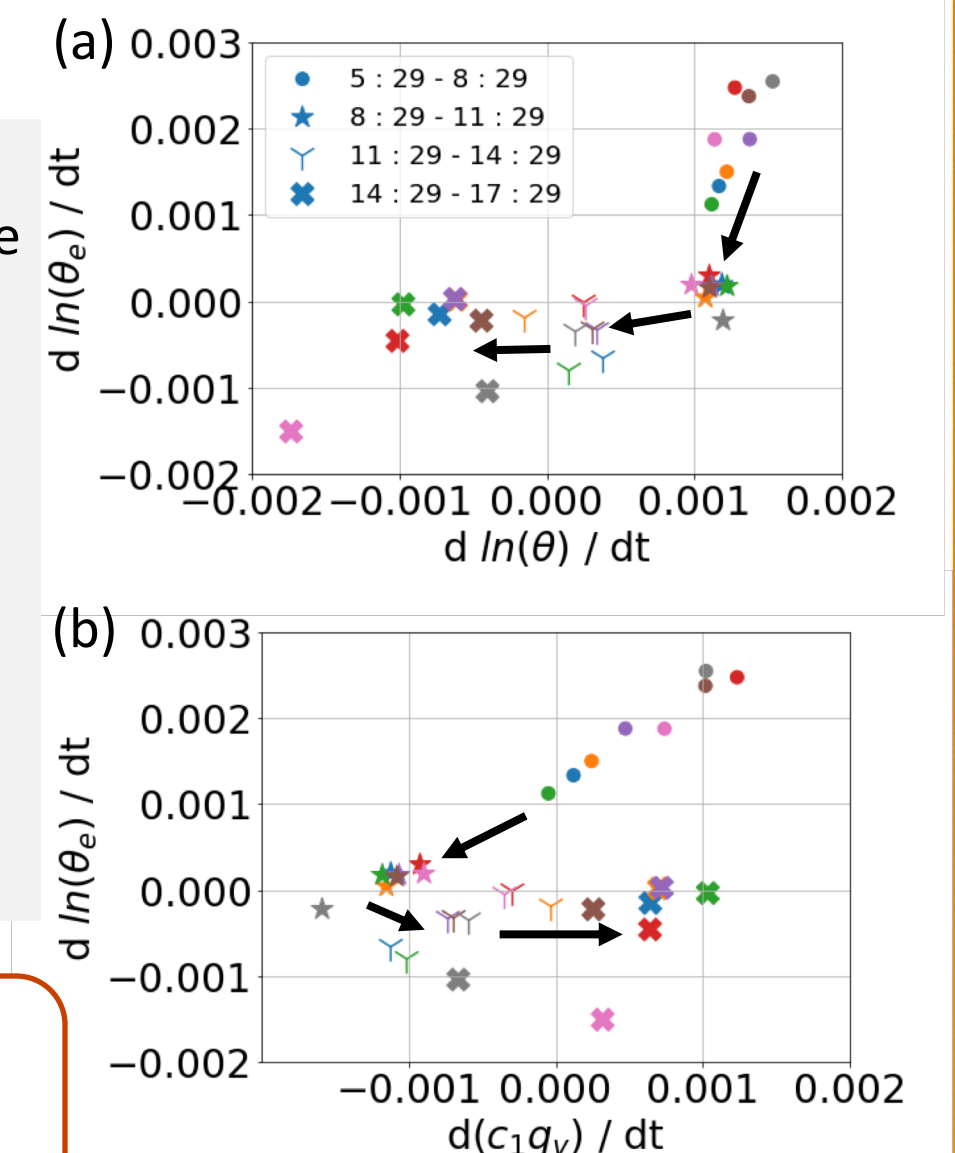


Figure 5. Contributions from (a) $d \ln \theta / dt$ and (b) $c_1 dq_v / dt$ to $d \ln \theta_e / dt$ for the 8 clusters from Figure 4.

Summary

- Convection points start with high θ_e , which is transported upward. As rain is formed, rain falls over the low θ_e environment.
- Convection prefers to occur over grassland and forest, rather cropland, urban and water.
- The spatial variations of q_v tendency are larger than those of θ tendency, and hence contribute more to the different trends of θ_e time series.

f_i : fraction of the i th land use type over domain (in Eq.3)

Notations:

θ_e : equivalent potential temperature
 θ_e' : θ_e spatial anomalies
 θ_{es}' : lowest level θ_e' (about 10m)

θ : potential temperature
 q_v : water vapor mixing ratio
 ρ : air density (in Eq.1)
 L_x or L_y : domain size (in Eq.1 and Eq.3)

LU: index of the land use type (in Eq.3)
 z : the height above the ground (in Eq.3)
 θ_{e0} : the θ_e of each isentropic contour (in Eq. 1 and Eq.3)
 $\delta()$: Dirac’s delta function (in Eq.1 and Eq.3)

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