

# Lecture 22

Continue with evolution of Earth's atmosphere  
(oxidation and reduction reactions on blackboard)

Start on “Snowball Earth”

# Evidence for an oxic transition

## Before the oxic transition: Reduced minerals (before 2.4 Ga)

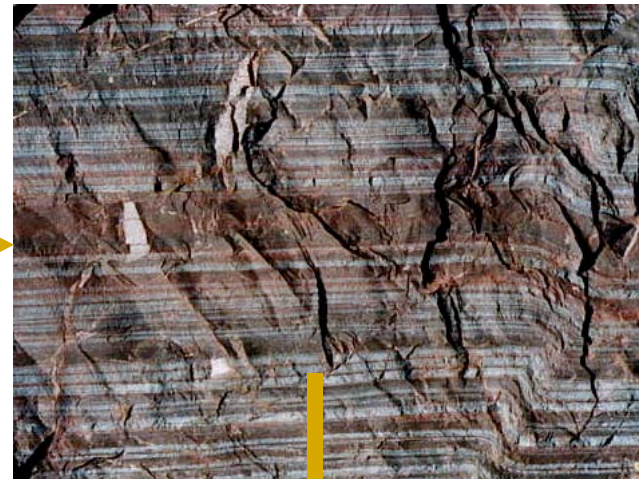
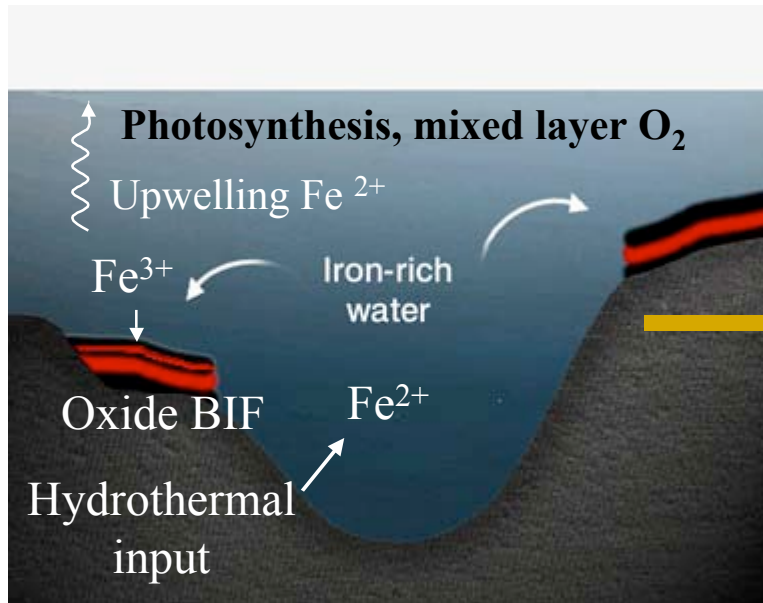
1. Banded Iron Formations.
2. Detrital Reduced minerals in riverbeds.
3. Paleosols (ancient soils) show Fe loss.

## After the transition: Oxidized minerals.

1. Red beds.
2. Decline of reduced indicators (BIFs, detrital) paleosols, Fe retained.

(Isotopic indicators are ‘smoking gun’ evidence (e.g. sulfur isotopes described in the textbook, p.217-218). But the chemistry is complex and we will skip this).

# Banded Iron Formations (BIFs)



Laminated chemical sediments  $\geq 15$  wt% Fe.

Japanese cars



Photo: A. Knoll, Harvard

# Litte change in the O<sub>2</sub> source

Remember that net O<sub>2</sub> input into the “atmosphere-ocean-surface” system comes from the burial of organic carbon.

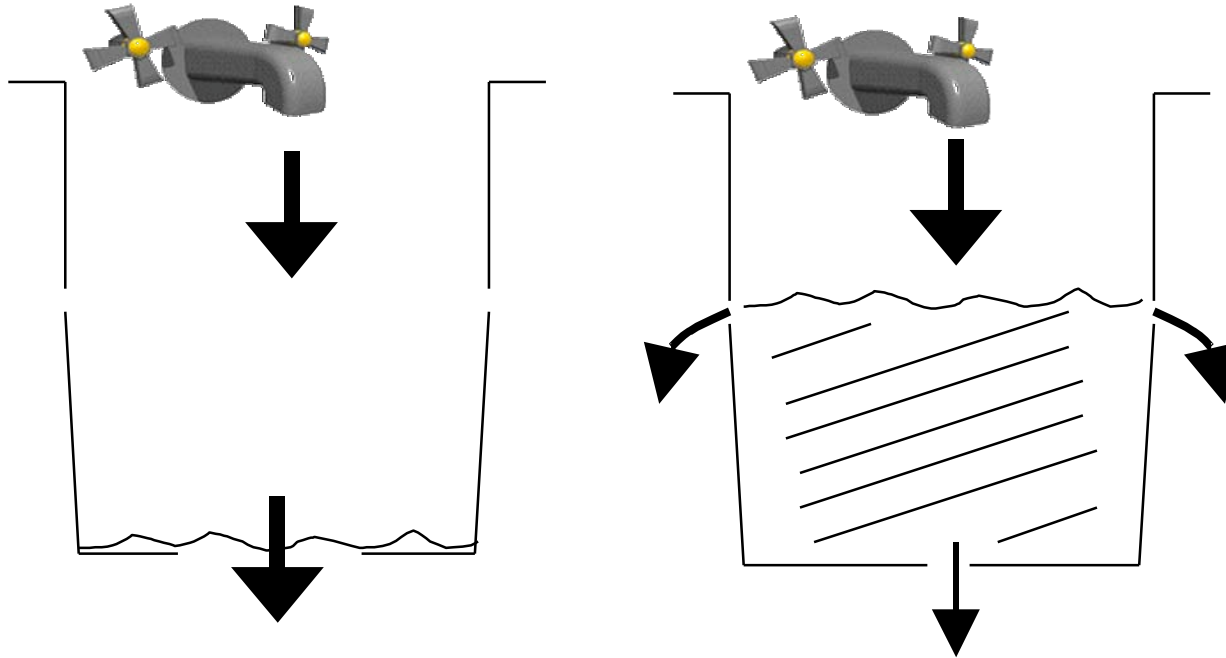
When we look at average sedimentary rock from 3.8 Ga to present, apart from a few relatively brief periods, remarkably we find:

- 1) roughly constant 0.5% by weight organic carbon (no change)
- 2) Isotopic evidence for litte change in organic carbon burial rates

If the source of O<sub>2</sub> has not changed, then what else could have caused a rise of O<sub>2</sub>?

Answer: the sink (i.e. mechanism for removal of O<sub>2</sub>).

## Leaky bucket analogy for growth in O<sub>2</sub> atmospheric levels



In each case: **flux in = total flux out**, but the water level is different.

Water level

- analogous to the O<sub>2</sub> amount in the atmosphere

Hole size in base

- analogous to amount of volcanic & metamorphic reducing gases (H<sub>2</sub>, CO) that react rapidly with O<sub>2</sub>

Holes in sides

- analogous to oxidation of continents (e.g. loss to red beds) that kicks in only at higher O<sub>2</sub> amounts.

# Theory for the rise of O<sub>2</sub>

It is thought that the Earth's rocky crust (and mantle) became more oxidized because of the loss of hydrogen to outer space.

A more oxidized crust and mantle produces more oxidized outgassed gases, and the amount of outgassed oxygen-consuming gases ( H<sub>2</sub>, CO, etc.) diminish. Implies more O<sub>2</sub>.

The set of reactions in the atmosphere that remove O<sub>2</sub> are highly **nonlinear**. We do not get a steady increase in O<sub>2</sub> with a steady decrease of reducing gases. Instead, O<sub>2</sub> reaches a critical point after which it leaps up in abundance.

# Aside: H<sub>2</sub> emissions in today's environment

REPORTS

Science, v. 300, p.1740 (2003)

## Potential Environmental Impact of a Hydrogen Economy on the Stratosphere

Tracey K. Tromp,<sup>1</sup> Run-Lie Shia,<sup>1</sup> Mark Allen,<sup>2</sup> John M. Eiler,<sup>1</sup>  
Y. L. Yung<sup>1\*</sup>

The widespread use of hydrogen fuel cells could have hitherto unknown environmental impacts due to unintended emissions of molecular hydrogen, including an increase in the abundance of water vapor in the stratosphere (plausibly by as much as ~1 part per million by volume). This would cause stratospheric cooling, enhancement of the heterogeneous chemistry that destroys ozone, an increase in noctilucent clouds, and changes in tropospheric chemistry and atmosphere-biosphere interactions.

2000 Presidential election:  
George W. Bush mocks Al Gore as an environmental extremist for wanting to eradicate the combustion engine in the future

January 2002:  
Pres. George W. Bush announces a \$1.2b program to research hydrogen fuel ....to eradicate the combustion engine in the future

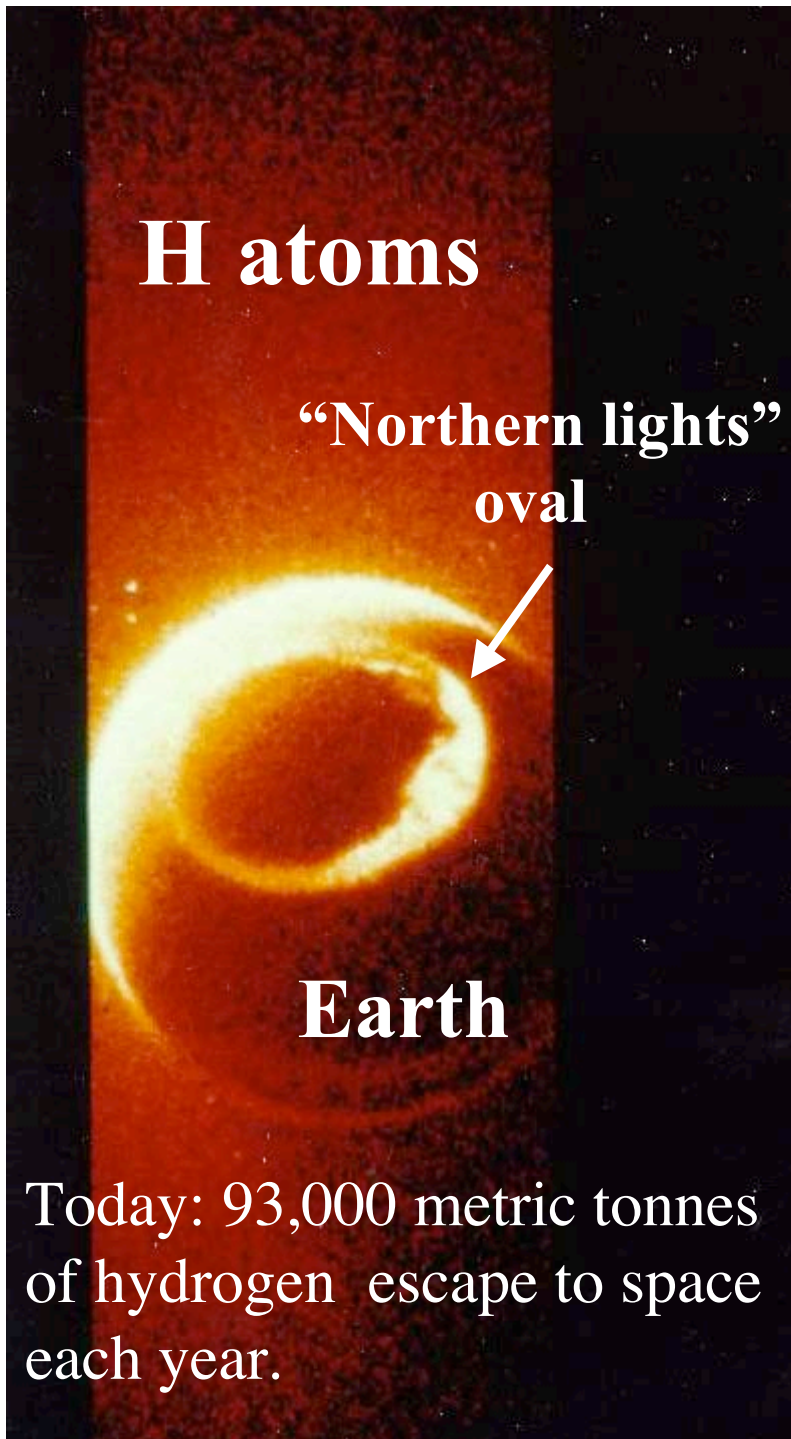
However, this paper took an extreme “worst case” view that 10% of hydrogen would leak from the future production/fueling system. This is probably a huge overestimate; more like 1%.

## Back to the Archean Earth: Enhanced H escape via methane, CH<sub>4</sub>

Today vast amounts of methane are consumed in reaction with oxygen in the atmosphere, leaving CH<sub>4</sub> at a trace abundance of only 1.7 ppmv.

In the low O<sub>2</sub> Archean atmosphere, biogenic methane would be abundant, at about ~1000 ppmv.

But what ultimately happens to such methane in the upper atmosphere? Answer: It gets photolyzed and hydrogen escapes away.



H atom **geocorona in red:**  
UV (121 nm) image of Earth.

- Today, half of these H atoms originate from microbial  $\text{CH}_4$  ; the other half from  $\text{H}_2\text{O}$  that makes it into the stratosphere.
  - With greater  $\text{CH}_4$ , H escape would be significant and oxidize the Earth
  - In the low- $\text{O}_2$  Archean,  $\text{CH}_4$  would be  $\sim 1000$  ppmv (compare 1.7 ppmv today).
- H escape rates were few hundred times greater.**

Methane's H escape causes O<sub>2</sub> gain



Photosynthesis + methanogenesis:

